-30C = 25808'

·200 = 21697'

.6 km

BASIC APPLICATION OF THE SKEW-T AND HODOGRAPH

IN SEVERE WEATHER FORECASTING 7 = 12354

,3′ km

I. THE SKEW-T

U — Thi

~5 FC (357m)

KYLE J GILLETT



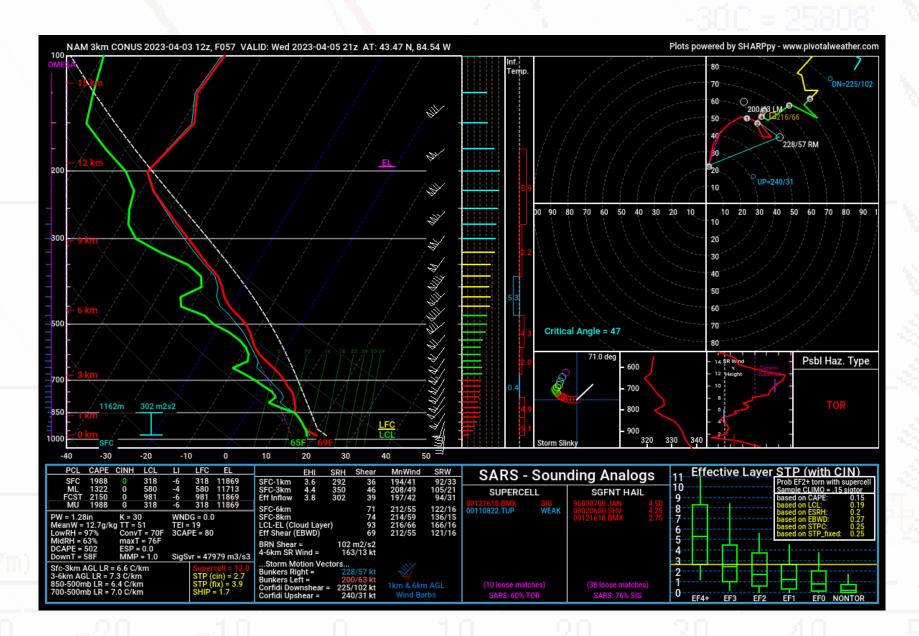


S • 12354'

SHEAR. LIFT. INSTABILITY. MOSITURE.

- 5 FC (357m)

WHAT'S A SOUNDING?

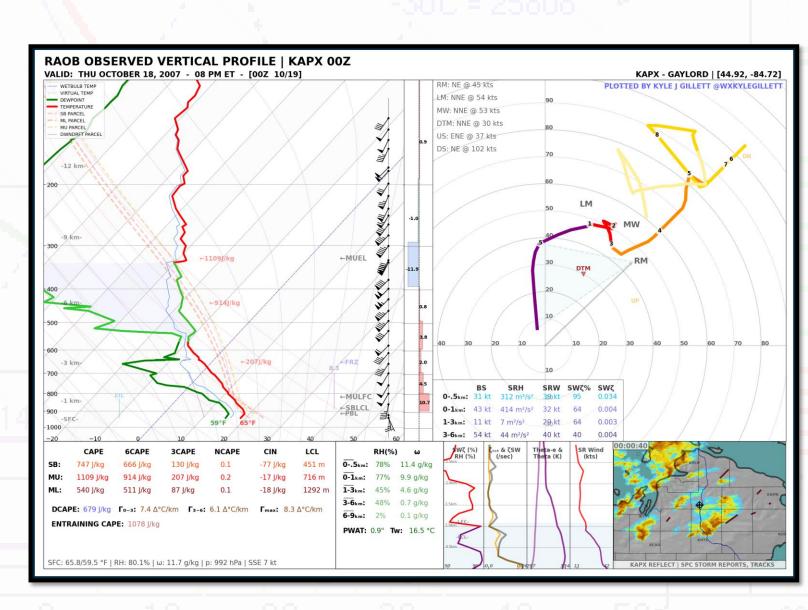


WHAT'S A SOUNDING?

TEMPERATURE DEWPOINT WIND

**That's *it*! **

Everything else plotted on a sounding is derived from those three values

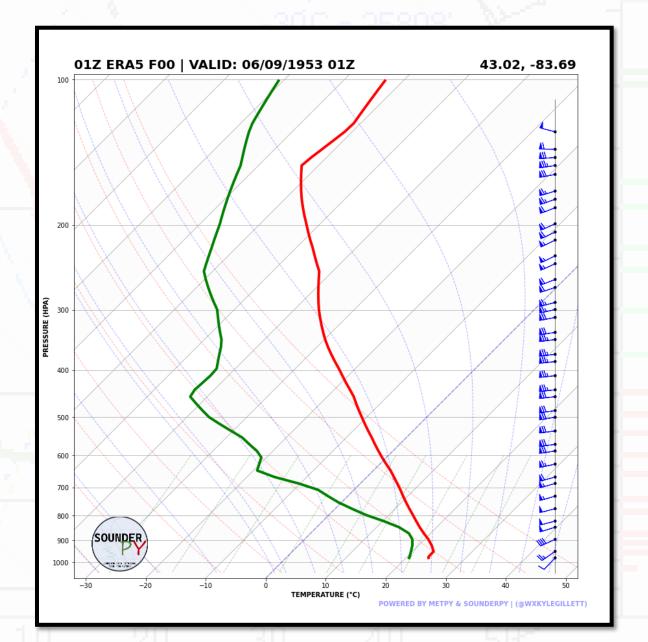


BASICS TO READING A SKEW-T

TEMPERATURE DEWPOINT WIND

THUS, WE CAN BEGIN TO ANWSER QUESTIONS ABOUT...

- INSTABILITY.
- MOSITURE.
- SHEAR.

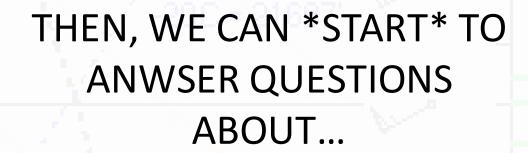


WHY THE SKEW-T IS IMPORTANT

IF WE CAN ANWSER QUESTIONS ABOUT...



- MOSITURE.
- SHEAR.





- STORM ORGANIZATION
- STORM MODE
- STORM THREATS
- AND LOTS MORE...



INSTABILITY: REVIEW

IN THIS SENSE, WE ARE CONSIDERING VERTICAL INSTABILITY

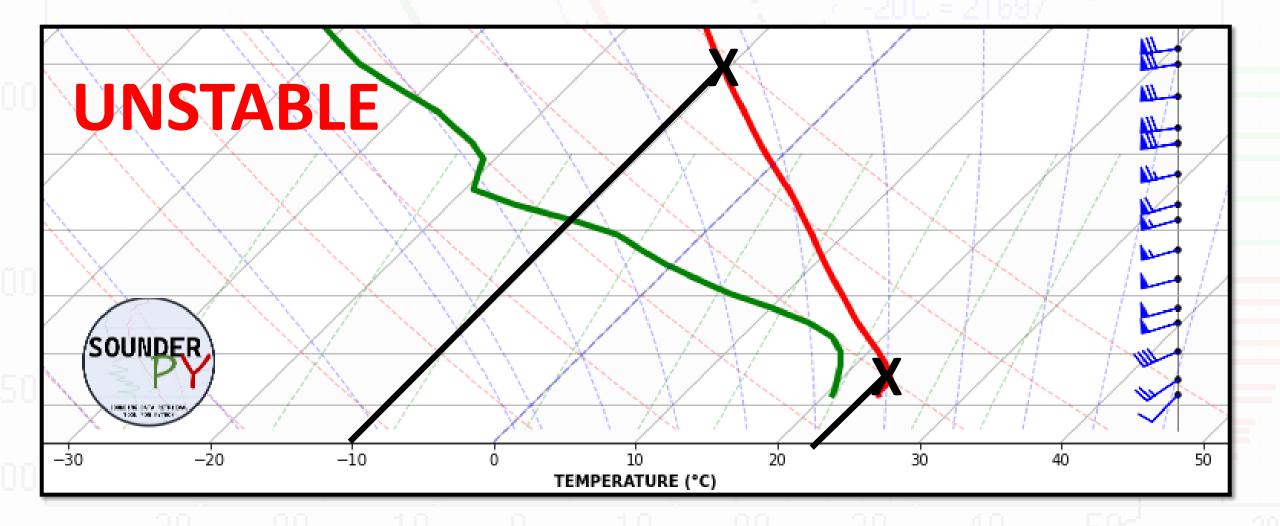
VERY SIMPLY: WE ARE LOOKING AT WHERE AIR IS RISING OR SINKING

ALSO VERY SIMPLY: STORMS WANT RISING AIR, OR INSTABILITY

INSTABILITY



SFC: 22°C -> 500 hPa: -10°C

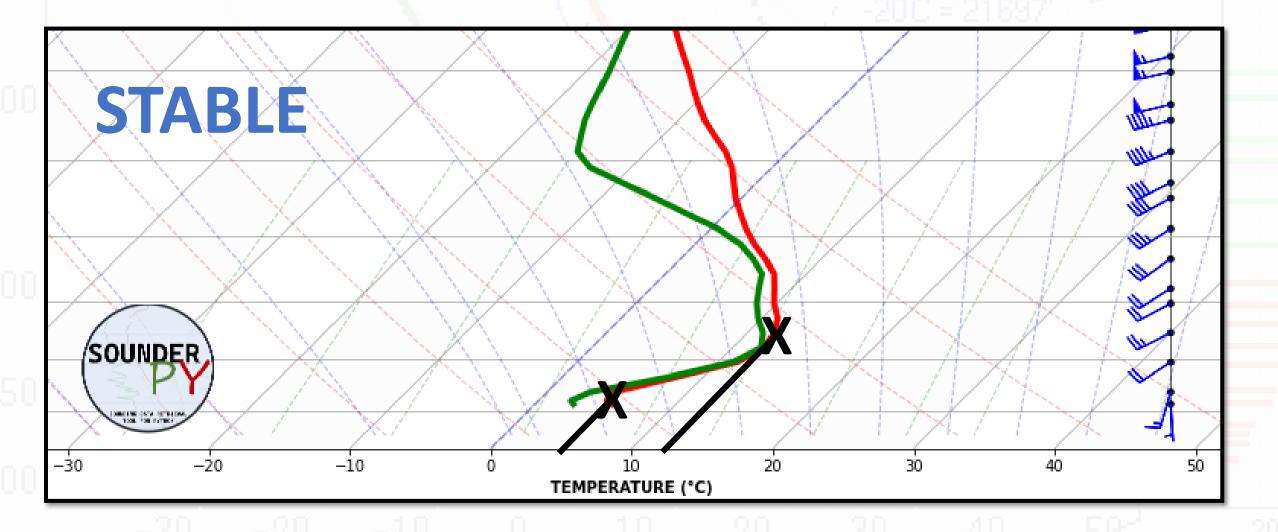


INSTABILITY

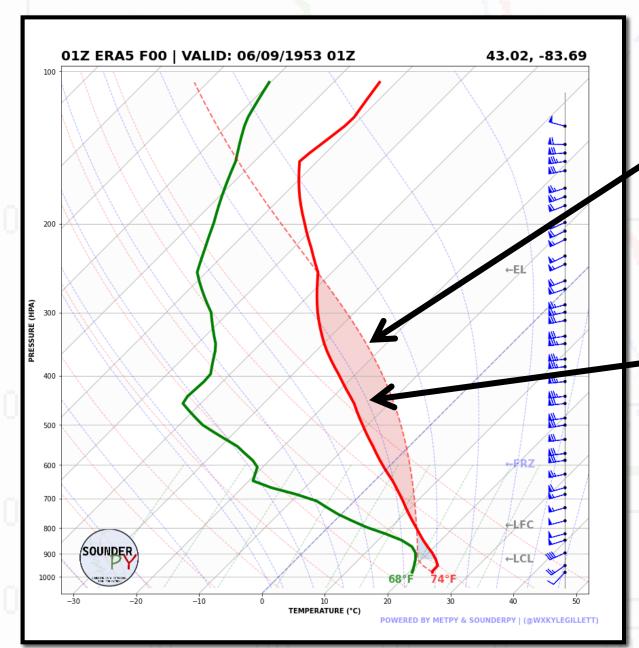
SFC: 5°C



950 hPa: 12°C



ASSESSING PROFILE STABILITY

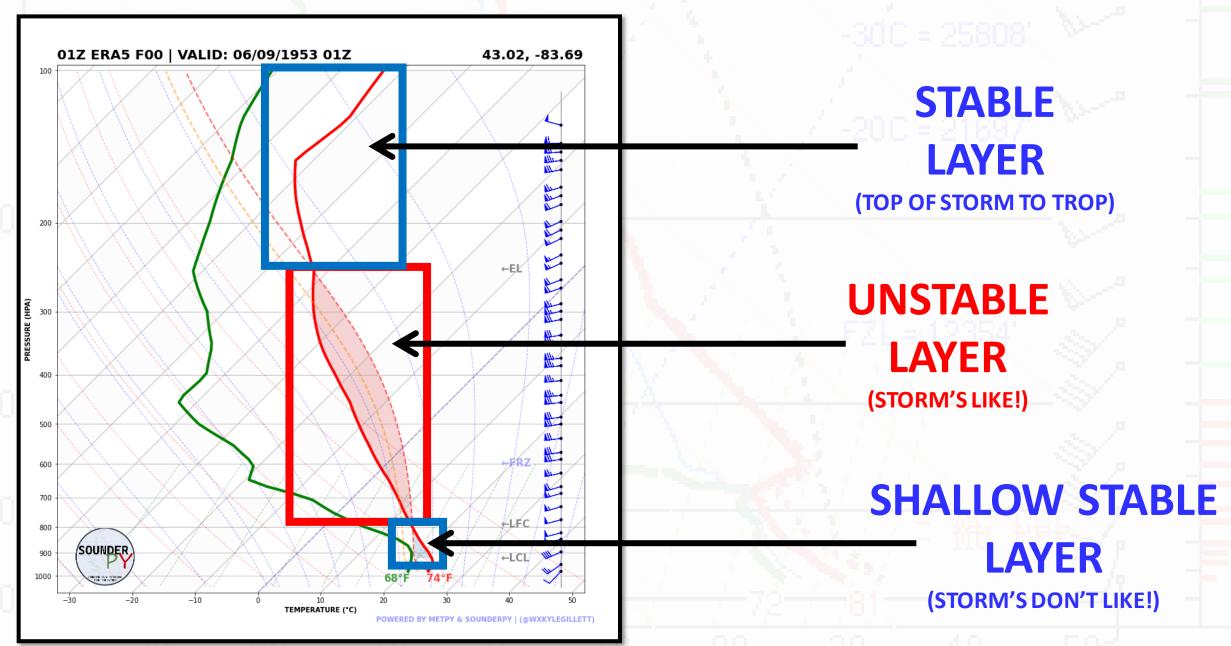


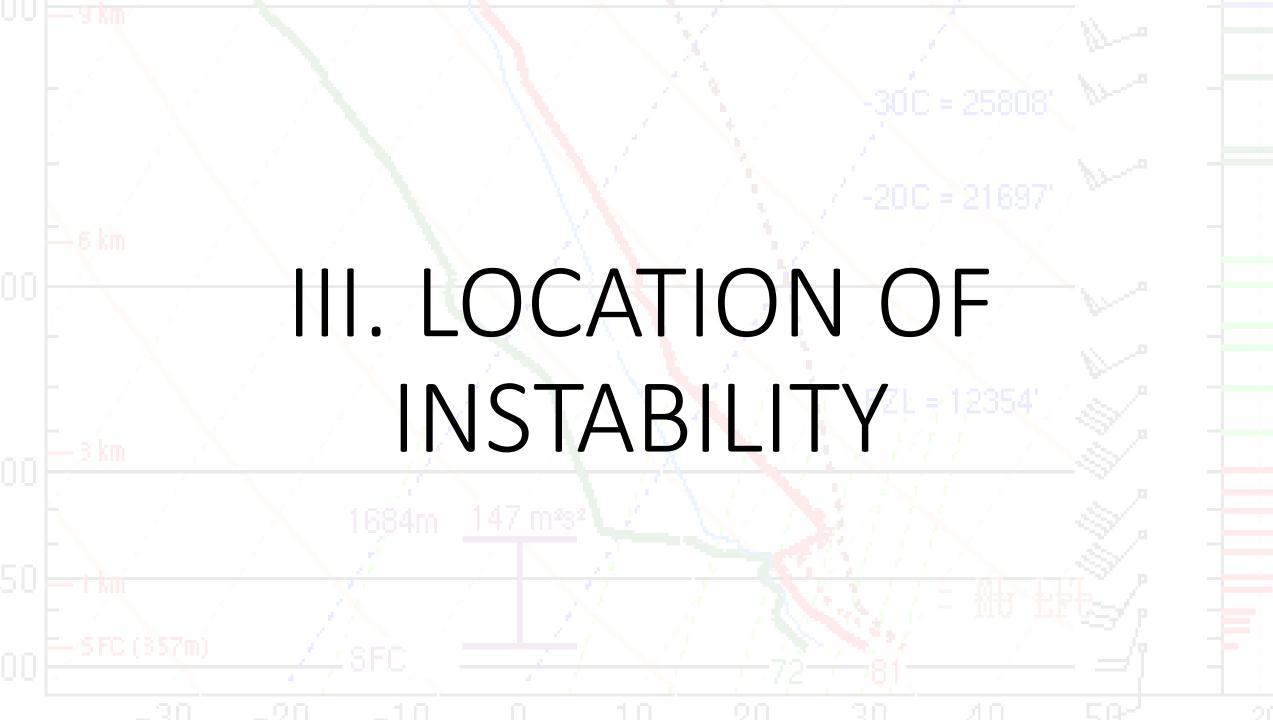
Parcels are unstable where the dashed 'parcel-trace' line is right of the environmental temperature line.

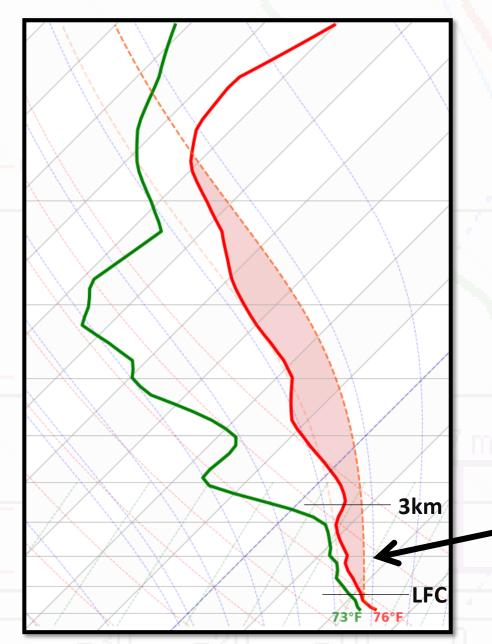
This is CAPE!

Conceptually, it's the area between a lifted parcel's temperature & the environmental temperature.

ASSESSING PROFILE STABILITY





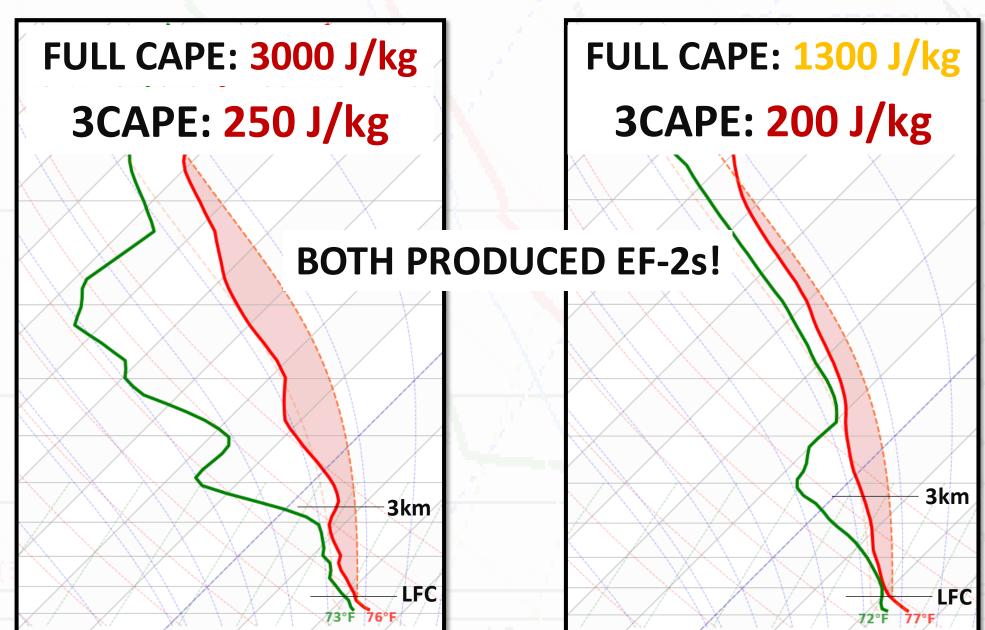


very important for tornadic storms.

Analytically, we can use **0-3km CAPE**to assess low-level instability

250 J/kg of 3CAPE!

FULL CAPE: 3000J/kg



3CAPE

VALUES

MODERATE

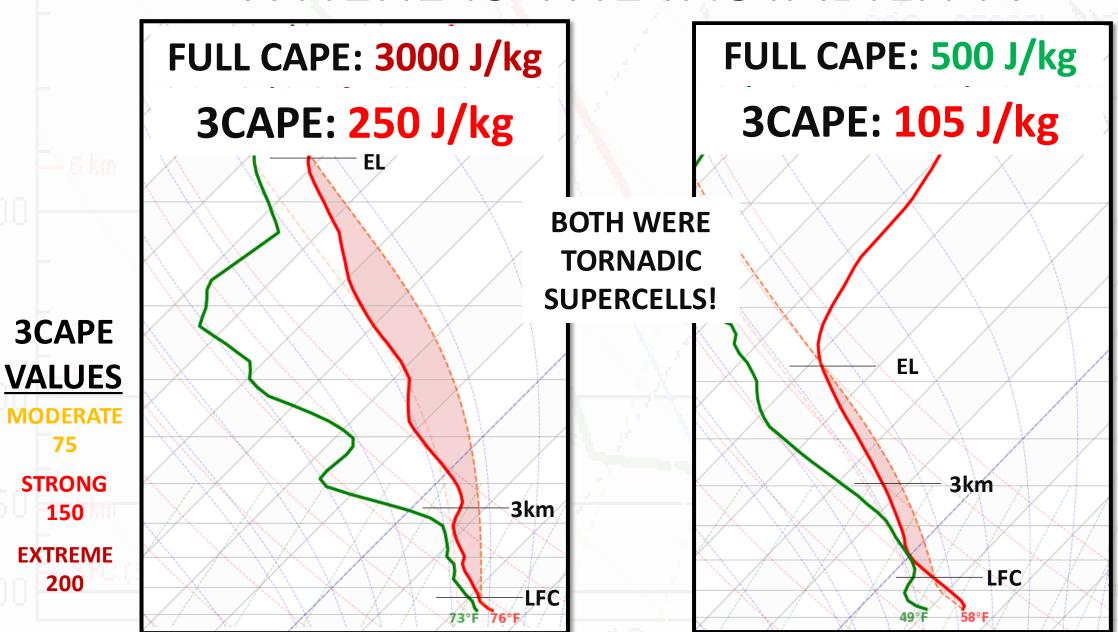
75

STRONG

150

EXTREME

200



3CAPE

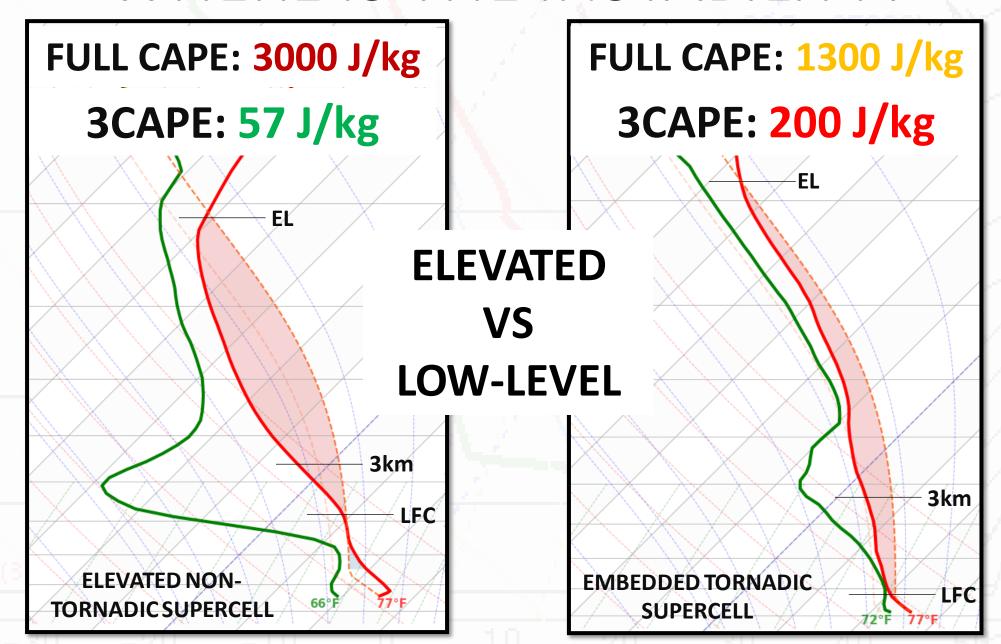
75

STRONG

150

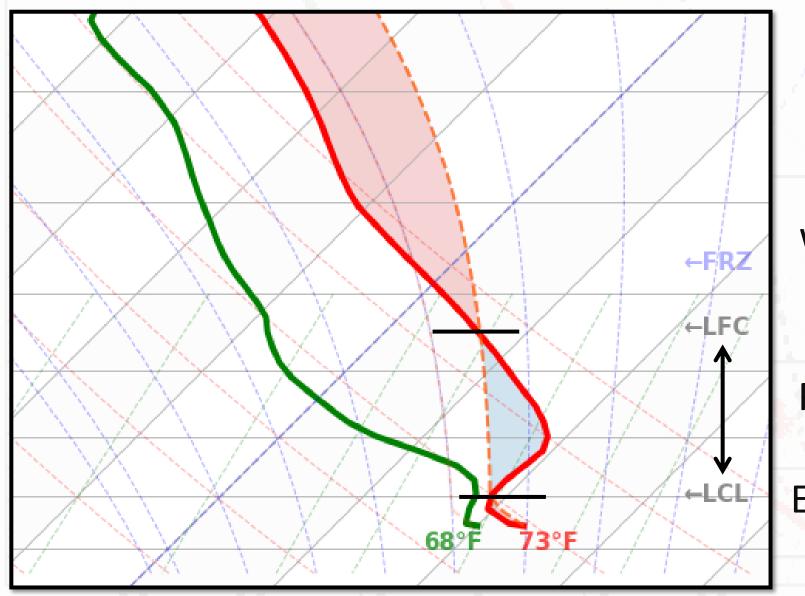
EXTREME

200





WHAT IS "THE CAP"?

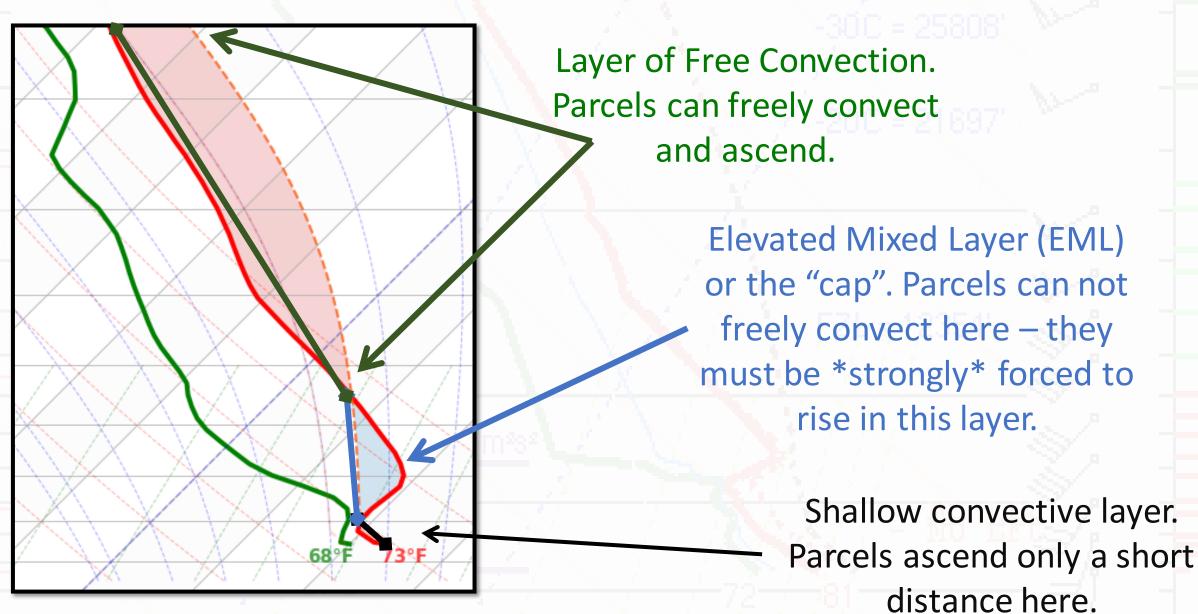


A STABLE LAYER IN THE LOW-MID LEVELS

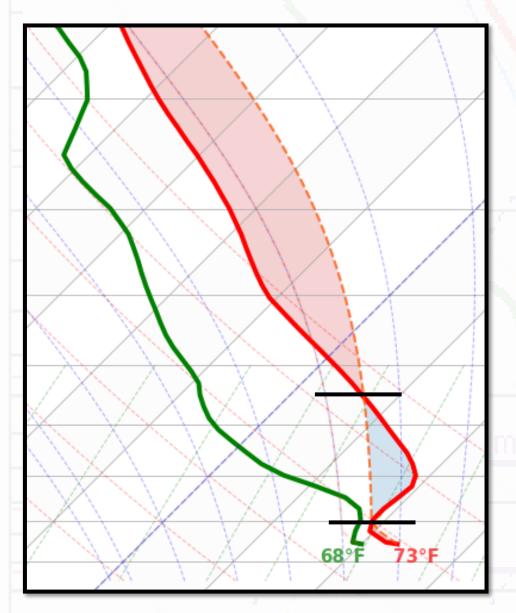
A WARM AND DRY
WEDGE OF AIR ABV THE
SURFACE

PARCELS ARE SUDDENLY
COOLER THAN THE
ENVIRONMENT, HALTING
CONVECTION

WHAT IS "THE CAP"?



WHAT IS "THE CAP"?



THE CAP <u>INHIBITS</u> DEEP MOIST SURFACE-BASED CONVECTION (THUS, WE FIND CIN OR CONVECTIVE INHIBITION -- LIKE INVERSE CAPE)

INSTABILITY IS BUILDING <u>ABOVE THE</u>

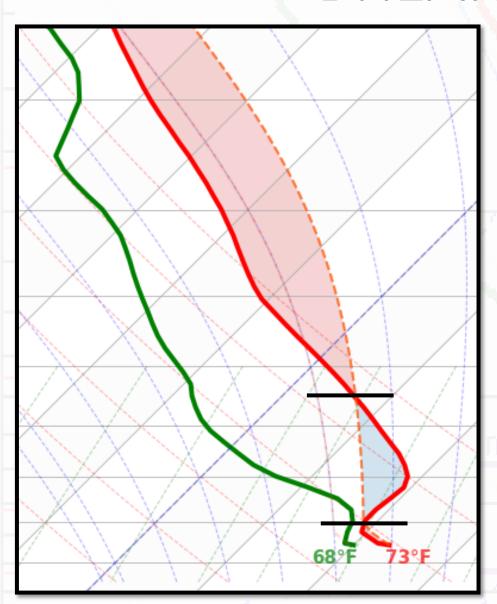
<u>CAP</u>, BUT STORMS CAN'T ACCESS IT

(UNLESS THEY ARE ELEVATED)

THE CAP PREVENTS SUFFICIENT CONVECTION FOR STORMS TO DEVELOP*

^{*} Sometimes a strong steady-state storm can *move into* a capped environment & remain strong – the updraft is *forcing* convection through the cap!

'BREAKING THE CAP'

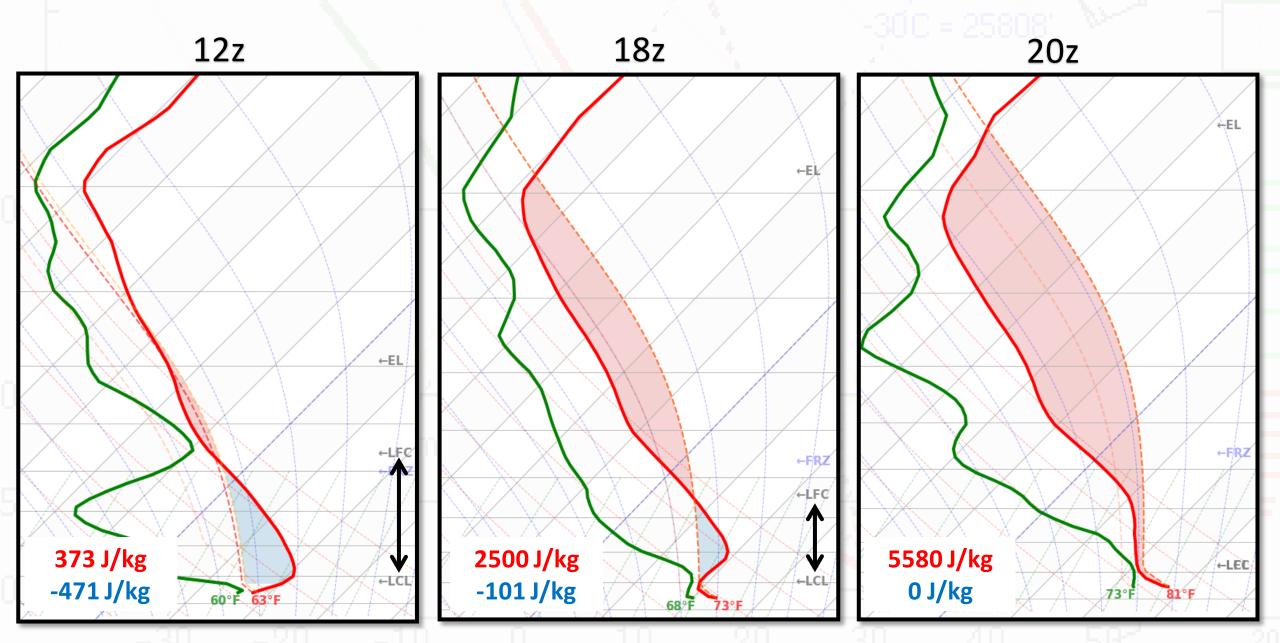


THE CAP CAN BE 'BROKEN' BY SUFFICIENT FORCING FROM...

A BOUNDARY (FRONT, DRYLINE)
LARGE SCALE ACSENT (SHORTWAVE,
GRAVITY WAVES)

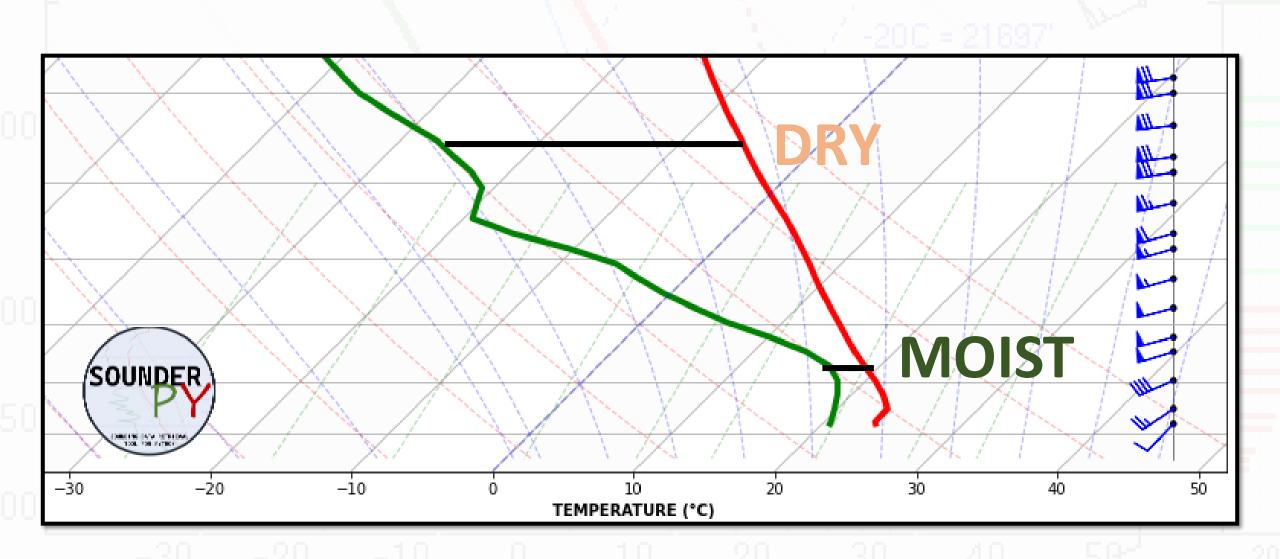
OR BY SUFFICIENT DIURNAL HEATING + MOISTURE TRANSPORT + MIXING.

THE "LOADED GUN" SOUNDING





ASSESSING THE MOISTURE PROFILE



ASSESSING THE MOISTURE PROFILE

$$CAPE = \int_{LFC}^{EL} \frac{T_v' - \overline{T_v}}{\overline{T_v}} dz$$

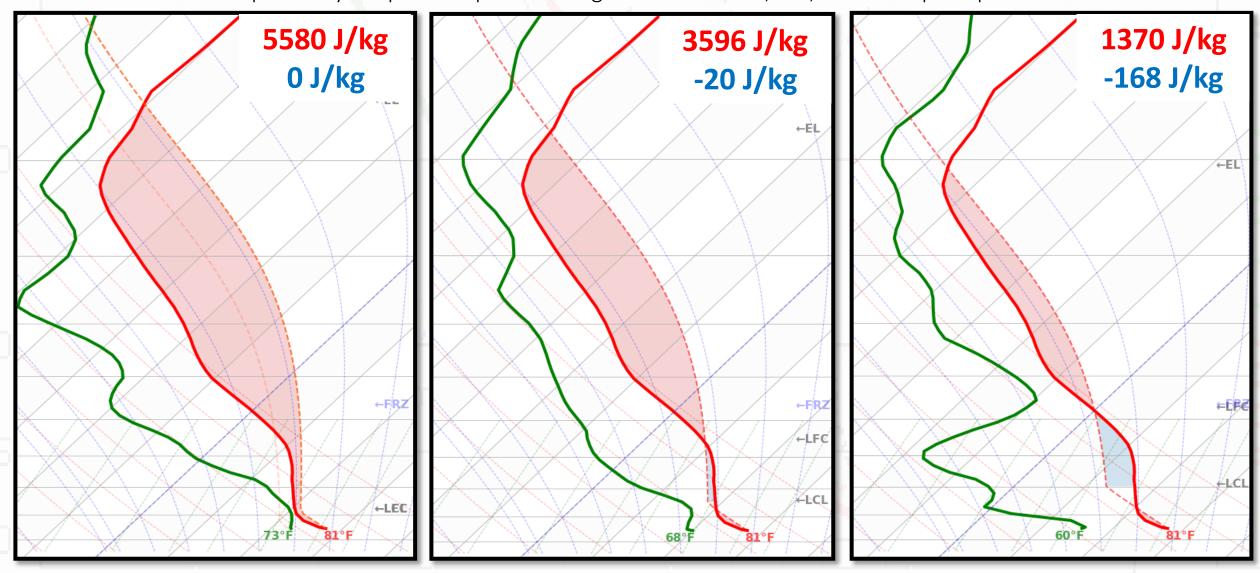
Virtual Temperature of a parcel

$$T_{v}' = T'(1 + 0.61q')$$

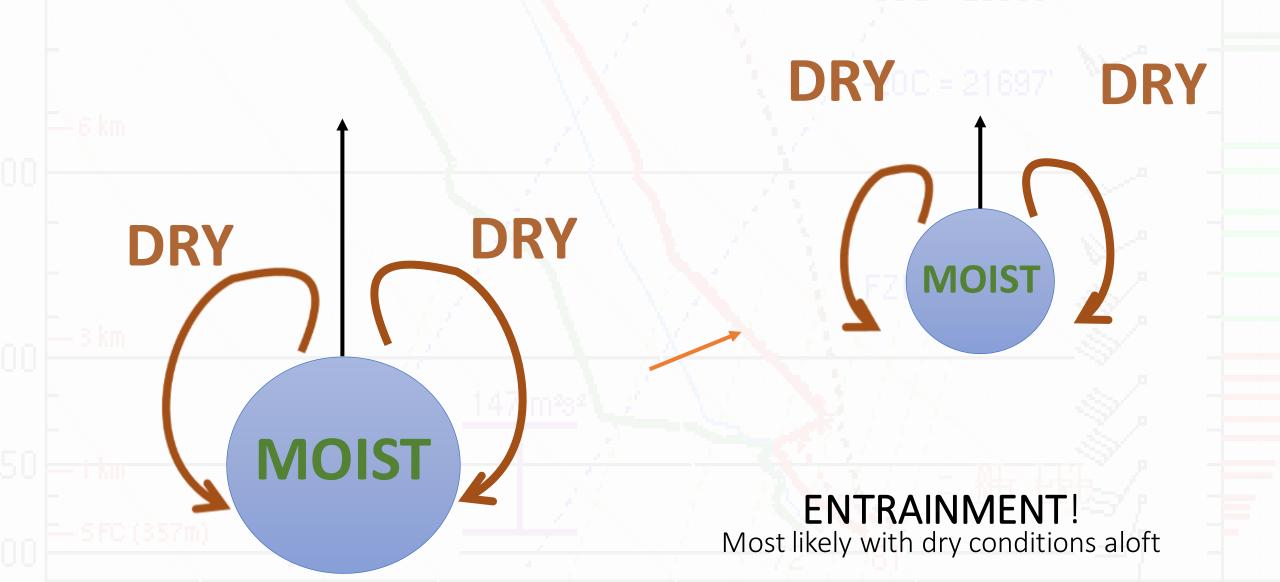
Specific Humidity = MOISTURE!

MOISTURE IMPACTS ON INSTABILITY

20z proximity temperature profile to Pilger 2014 with 20z, 18z, & 12z dewpoint profiles



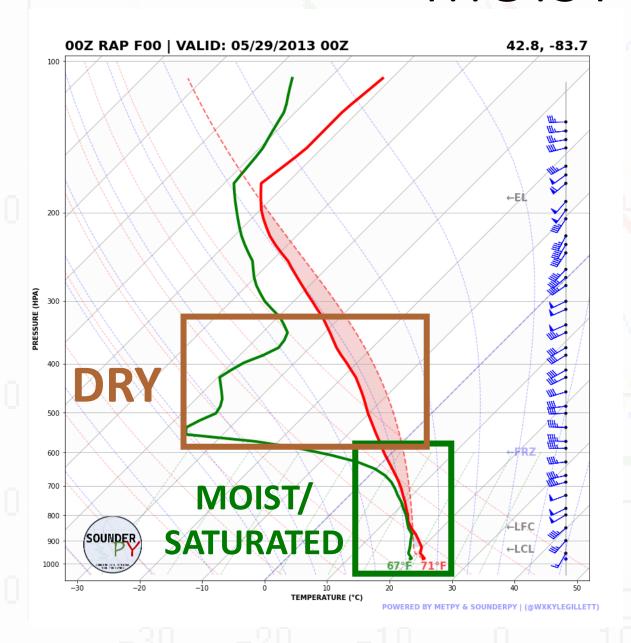
DILUTION OF CAPE

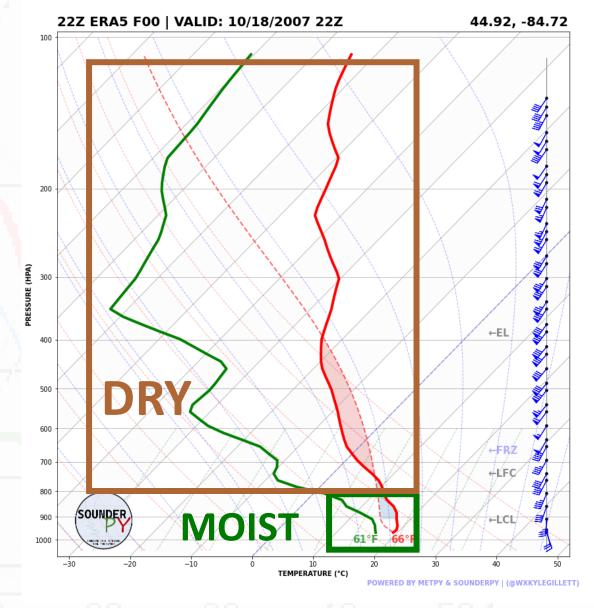


-30C = 25808' -20C = 21697'

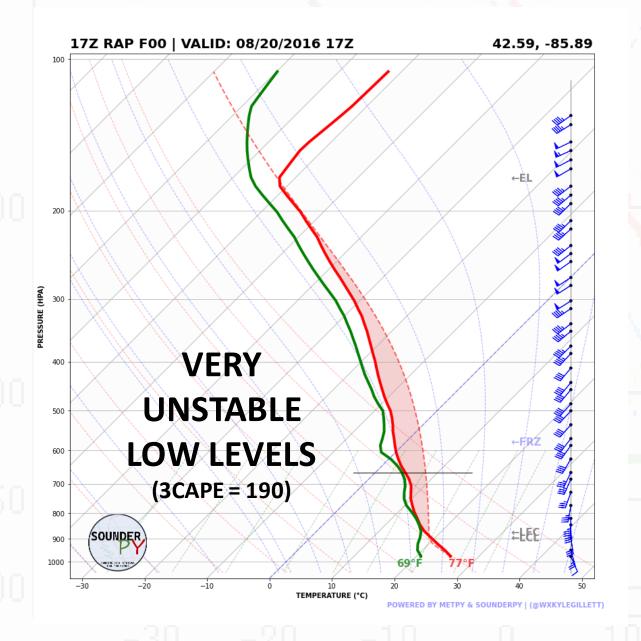
VI. EXAMPLE SEVERE CONVECTION SKEW-Ts

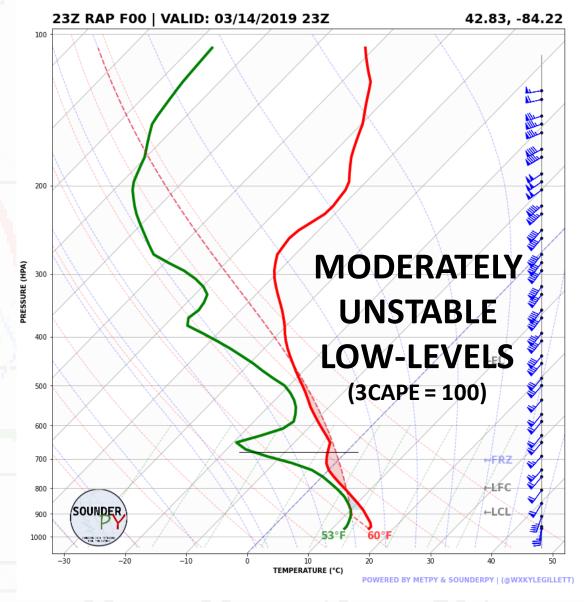
MOIST vs DRY



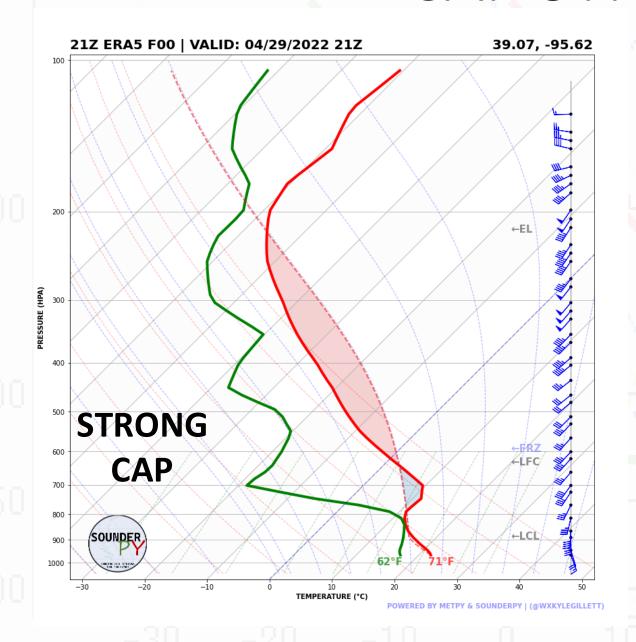


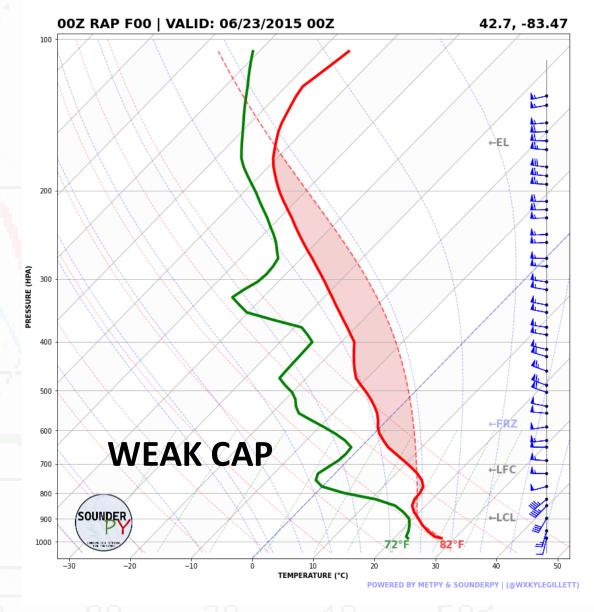
LOW LEVEL INSTABILITY





CAP STRENGTH





THAT'S THE BASICS OF THE SKEW-T ITSELF.

THE HODOGRAPH IS NEEDED TO PROVIDE FURTHER INSIGHT.

THAT'S NEXT TIME;)